Radiative Transfer Theory to Estimate Medium Characteristics

Measure

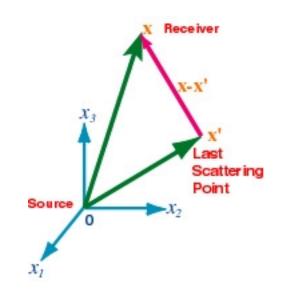
scattering coefficient, intrinsic attenuation

Terminology:

Q_i-1 Intrinsic Attenuation

Q_s⁻¹ Scattering Attenuation g Scattering Coefficient

$$g_0 \equiv \frac{1}{4\pi} \oint g d\Omega = \ell^{-1} = Q_s^{-1} k$$



Albedo: ratio of scattering to total attenuation

$$B_0 \equiv \frac{Q_S^{-1}}{Q_i^{-1} + Q_S^{-1}} = \frac{g_0 \beta_0}{g_0 \beta_0 + Q_i^{-1} \omega}$$

Terminology and Relationships:

- Q_i-1 Intrinsic Attenuation
- Q_s⁻¹ Scattering Attenuation

$$Q_T^{-1} = Q_S^{-1} + Q_i^{-1}$$

$$g_0 \equiv \frac{1}{4\pi} \oint g d\Omega = \ell^{-1} = Q_s^{-1} k$$

$$B_0 \equiv \frac{Q_S^{-1}}{Q_i^{-1} + Q_S^{-1}} = \frac{g_0 \beta_0}{g_0 \beta_0 + Q_i^{-1} \omega}$$

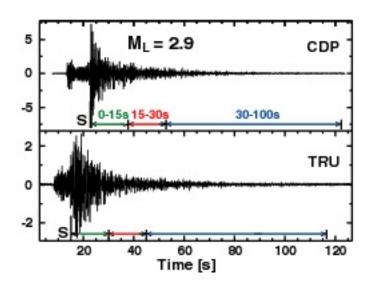
First measurements done by Wu and Aki (1988)

- Formulated radiative transfer theory for seismology
- Steady-state solution (Ishimaru, 1978)
- Measured wavefield energy vs. distance
 - compared observations with theory
 - Found scattering strength extremely strong $Q_s^{-1} >> Q_i^{-1}$

Multiple Lapse-Time Window Analysis Method

- Based on time-domain solution of radiative transfer theory equation
- Analyze Integrated
 Energy Density vs time
 and space
- Allows Separation of Scattering and Intrinsic Attenuation

Integrate Energy in Windows whose Times are Referenced to S-wave Arrival Time



Procedure

 Define 3 time windows and integrate energy in these windows

$$EI_{1}(f)_{kj} = \rho_{0} \int_{0}^{15s} |\dot{u}_{kj}^{S}(t;f)|^{2} dt,$$

$$EI_{2}(f)_{kj} = \rho_{0} \int_{15s}^{30s} |\dot{u}_{kj}^{S}(t;f)|^{2} dt,$$

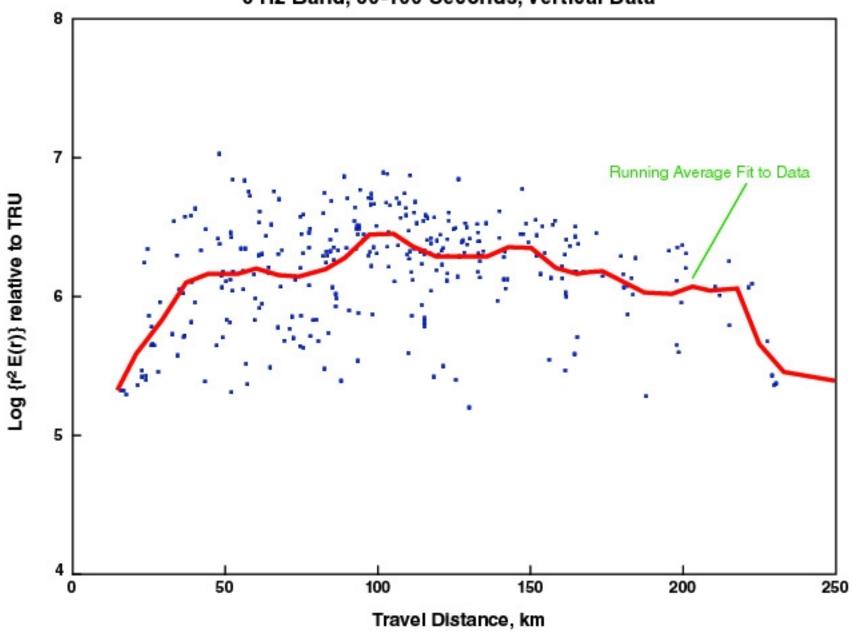
$$EI_{3}(f)_{kj} = \rho_{0} \int_{30s}^{100s} |\dot{u}_{kj}^{S}(t;f)|^{2} dt$$

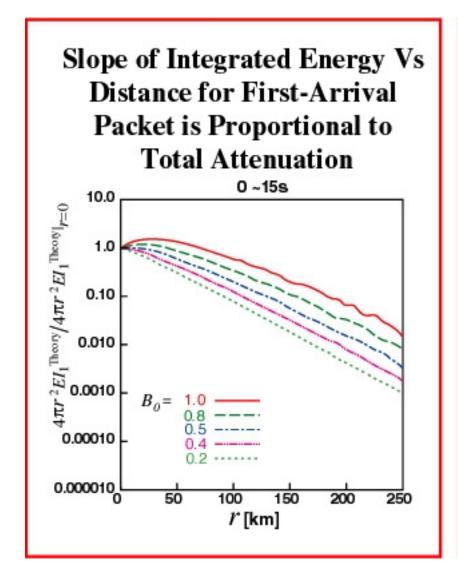
• Use Coda Normalization to correct data for station amplification and source size

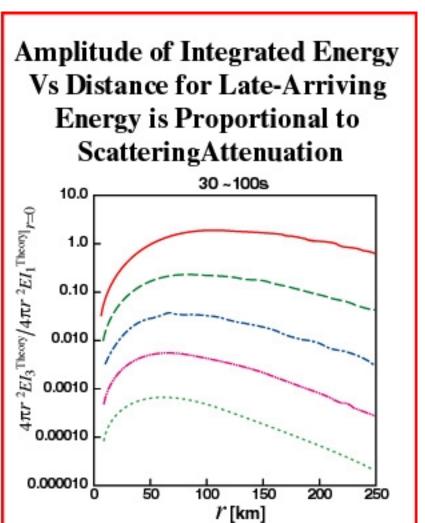
Procedure #1

- Compare running average fits of data to model
- Comparing fits for three (or more?) time windows gives reasonably unique fit
- Could define "best" fit using some curve fitting procedure

3 Hz Band, 30-100 Seconds, Vertical Data



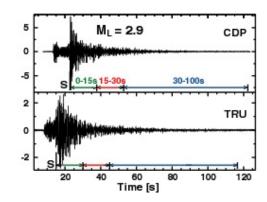




Integrate Energy in Windows whose Times are Referenced to S-wave Arrival Time

Procedure #2

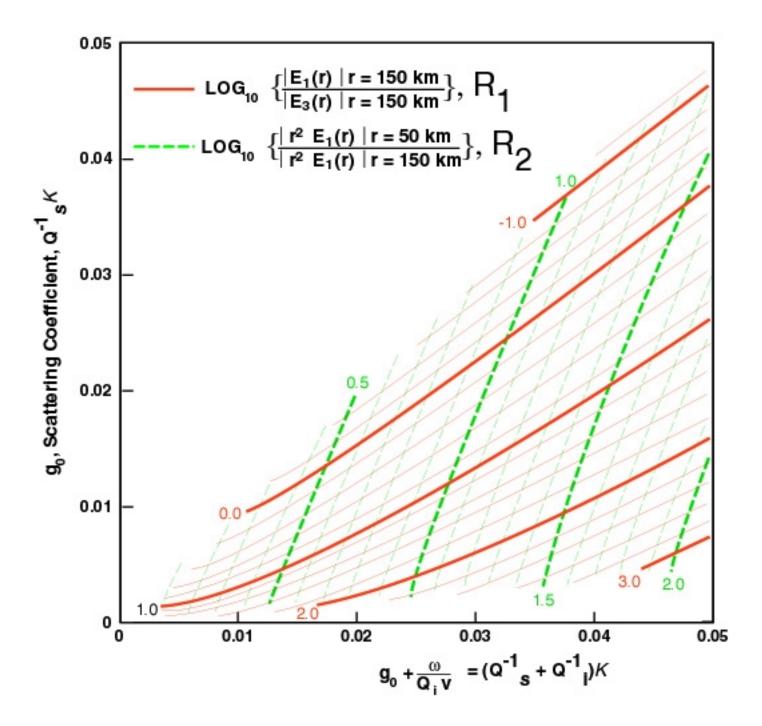
Measure following two



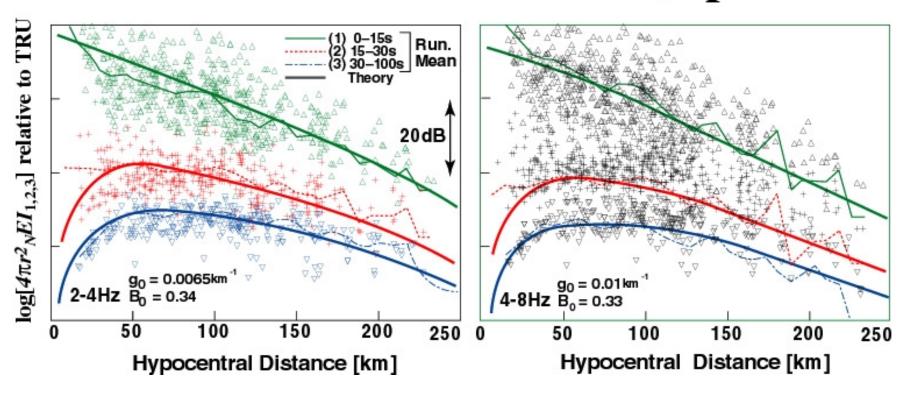
$$R_{1} = \log \left[\frac{\langle NEI_{1} \rangle_{D}|_{r=150 \text{ km}}}{\langle NEI_{3} \rangle_{D}|_{r=150 \text{ km}}} \right]$$

$$R_{1} = \log \left| \frac{\langle_{N}EI_{1}\rangle_{D}|_{r=150 \text{ km}}}{\langle_{N}EI_{3}\rangle_{D}|_{r=150 \text{ km}}} \right| \qquad R_{2} = \log \left| \frac{4\pi r^{2}\langle_{N}EI_{1}\rangle_{D}|_{r=50 \text{ km}}}{4\pi r^{2}\langle_{N}EI_{1}\rangle_{D}|_{r=150 \text{ km}}} \right|$$

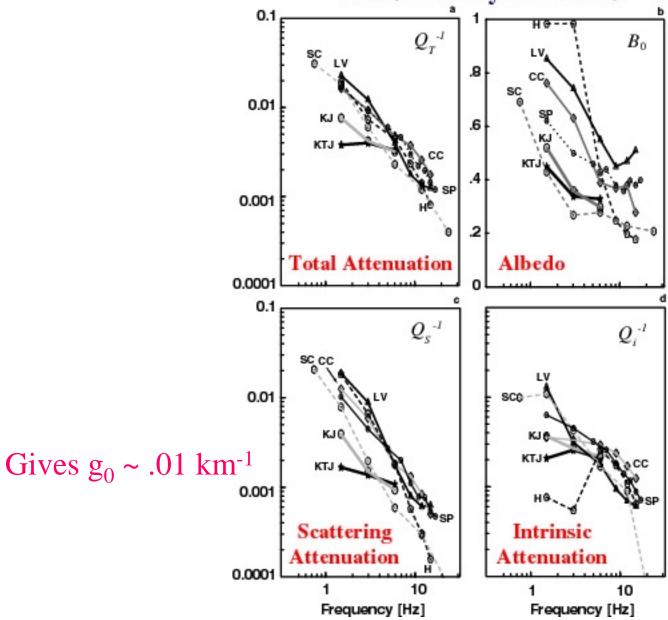
- R₁ is proportional to amount of scattering
 - Large R₁ means no scattered waves
- R₂ is proportional to total attenuation
 - Slope of amplitude vs distance is classic way to determine total attenuation



Fits to Data Collected in Japan



Results Obtained with Multiple Lapse-Time Window Analysis Method



- Feustel et al. (1996) studied mine seismicity (400-1600 Hz) and found $g_0 \sim .25 \text{ m}^{-1}$ which corresponds to characteristic length of mapped fractures
- Frequency dependence of g_0 ?

Other approaches for estimating medium characteristics:

- Fitting curves using optimization procedure
- Anisotropic scattering model

Conclusions

- Scattering models can be used to obtain useful information about the Earth
- Parameters vary with tectonic region and do provide information to characterize regions
- Coda normalization method works independently of any clear knowledge of its basis